## Supporting Information

## Origin of Magnetism in Hydrothermally Aged 2-line Ferrihydrite Suspensions

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Number of pages: 11 Number of Figures: 7 Number of Tables: 1 **XMCD spectrum.** An XMCD spectrum for 6-line ferrihydrite was measured at room temperature, which is above the blocking temperature of 6-line ferrihydrite. It is, therefore, expected that the signal will be close to the noise level of the synchrotron detectors.

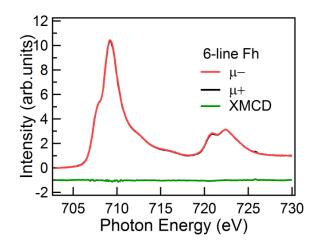
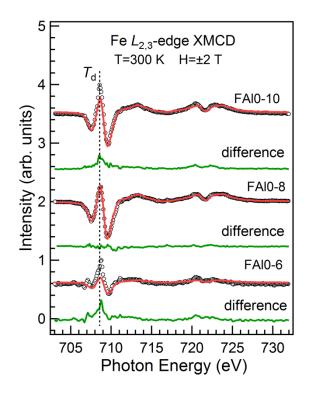


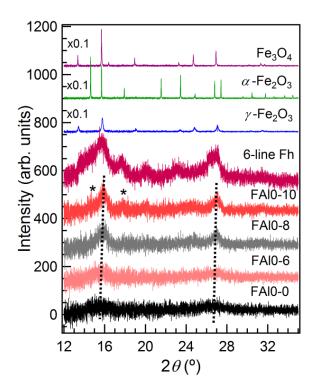
Figure S1. XMCD spectrum of 6-line ferrihydrite (Fh) measured at room temperature.

XMCD spectra of hydrothermally aged ferrihydrite are fitted by a linear combination of  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> and Fe<sub>3</sub>O<sub>4</sub> reference spectra (**Figure S2**). The existence of Fe<sub>3</sub>O<sub>4</sub> during hydrothermal aging can be ruled out because the fitted spectra have a contribution only from the  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> spectrum with an uncertainty no larger than 3%, which supports the conclusions drawn from Fe *L*-edge XAS and *K*-edge XANES curve-fitting results. The fitting curve for FAI0-08 almost perfectly reproduces the experimental data. However, XMCD spectra for 6 h and 10 h aged ferrihydrite have stronger positive tetrahedral Fe<sup>3+</sup> ( $T_d$ (Fe<sup>3+</sup>)) than that in the fitted spectra, which suggests that the  $T_d$ (Fe<sup>3+</sup>) could be stronger in those samples than in  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> possibly due to a complicated spin arrangement among different sites in the ferrihydrite core and the  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> shell structures. The XMCD spectra calculated as the intensity difference of XAS spectra naturally have a much larger noise to signal ratio, which inevitably also leads to a larger fitting uncertainty.



**Figure S2.** Fe  $L_{2,3}$ -edge XMCD spectra for the three studied 2-line ferrihydrite samples with different hydrothermal aging times. The red curves represent best fits by linear combination of  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> and Fe<sub>3</sub>O<sub>4</sub> reference spectra. Difference spectra between fitted and experimental data are shown in green.

**Crystal Structure.** Synchrotron-based X-ray diffraction (XRD) measurements were made at the BL14B1 beamline of the Shanghai Synchrotron Radiation Facility (SSRF) with a wavelength of 0.6887 Å.<sup>1</sup> Measurements were made in spinning capillary mode, with samples loaded into a 0.5-mm glass capillary. A Mythen 1 K linear detector, which is mounted on the diffraction arm ~76 cm away from the sample, was employed for data acquisition. The typical angular resolution is ~ 0.0037°/channel (1280 channels cover a  $2\theta$  range of 4.8°). Acquisition time for one scan is ~1 minute. The data were analyzed using the CMPR program1 for background subtraction (empty capillary).



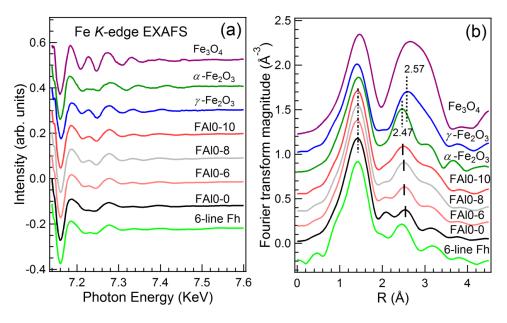
**Figure S3.** X-ray diffraction patterns for the four studied 2-line ferrihydrite samples (wavelength of 0.6887 Å). Standard diffraction patterns for 6-line ferrihydrite,  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>,  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> and Fe<sub>3</sub>O<sub>4</sub> are also shown.

The crystal structure of the reference samples and the studied 2-line ferrihydrite samples was further characterized by XRD (**Figure S3**). Two prominent diffraction peaks for samples aged for 10 h are located at ~15.86° and ~26.95°, respectively. The corresponding *d* spacings are 2.50 Å and 1.48 Å, respectively. The diffraction intensity increases with hydrothermal aging time, which demonstrates that crystallinity is enhanced during hydrothermal aging. All peaks shift to higher  $2\theta$  angles during hydrothermal aging, which suggests a systematic structural change in the samples. To explain this behavior, the diffraction pattern for the sample after 10 hours of aging at 450 K is compared with diffraction patterns for  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>,  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub>, Fe<sub>3</sub>O<sub>4</sub> and 6-line ferrihydrite standards. Although all of the standards have diffraction peaks at around 15.86° and 26.95°, the two peaks for Fe<sub>3</sub>O<sub>4</sub> occur at slightly smaller angles than for  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> and for the sample after 10 hours of aging at 450 K. Moreover, based on diffraction peaks at ~15.86° and ~26.95°, the calculated lattice parameter is not consistent with that of a cubic unit cell for Fe<sub>3</sub>O<sub>4</sub>. For a tetragonal structure, lattice parameters of a = b = 8.49 Å and c = 24.70 Å are obtained. Thus, Fe<sub>3</sub>O<sub>4</sub> should not exist in this sample. Moreover, the weak (asterisked) peaks in **Figure S3** may originate from 6-line ferrihydrite and/or  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>. The peak at ~20.09° after 10 hours of aging at 450 K is attributed to 6-line ferrihydrite. This sample should, therefore, contain both  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> and  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>, but not Fe<sub>3</sub>O<sub>4</sub>, which supports our XMCD and XAS results.

Local structure around Fe atoms: To understand the evolution of the local structure around Fe atoms in ferrihydrite with increasing hydrothermal aging time, *K*-edge extended X-ray absorption fine structure (EXAFS) spectra were measured for standard  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>,  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub>, Fe<sub>3</sub>O<sub>4</sub>, 6-line ferrihydrite and the four studied 2-line ferrihydrite samples (Figure S4(a)). Corresponding phase-shift uncorrected *K*-edge Fourier transform EXAFS spectra are shown in Figure S4(b). The first single coordination peak in the 1~2 Å region corresponds to the first Fe-O coordination shells in Fe atoms. The second peak corresponds to the Fe-Fe interatomic distance in second coordination shells. Peaks for  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>,  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> and Fe<sub>3</sub>O<sub>4</sub> standards are located at ~2.47 Å and ~2.57 Å, respectively.

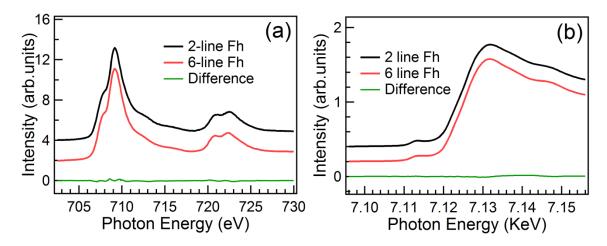
In general, EXAFS features resemble those of the  $(\alpha - /\gamma -)$ Fe<sub>2</sub>O<sub>3</sub> standards and are clearly different from those of the Fe<sub>3</sub>O<sub>4</sub> standard, which indicates that Fe<sub>3</sub>O<sub>4</sub> does not exist in these samples. The Fe-Fe coordination peak of 2-line ferrihydrite is located at ~2.57 Å, which is comparable with the value for  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub>. Its position decreases gradually to lower values with hydrothermal aging time, and reaches ~2.50 Å after 10 h of aging at 450 K, which suggests a systematic structural change, such as a reduced distortion due to an ordered structure or

dehydration of the mineral surface that resulted from decreased Fe-Fe interatomic distance. It is accepted that spectra for multi-site minerals indicate an intermediate structure. Peak positions for the studied ferrihydrite samples associated with Fe-O coordination (dashed-dot line in **Figure S4**(b)) and Fe-Fe coordination (solid bars in **Figure S4**(b)) all lie between values for  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub> and  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub>, which indicates that the aged samples contain both  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> and  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>, but not Fe<sub>3</sub>O<sub>4</sub>, which supports our XMCD and XAS results.



**Figure S4.** (a) *K*-edge EXAFS spectra for four studied 2-line ferrihydrite samples and standard samples of  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>,  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub>, Fe<sub>3</sub>O<sub>4</sub> and 6-line ferrihydrite throughout the continuum energy region and (b) corresponding phased-shift uncorrected *K*-edge Fourier transform EXAFS spectra. Fe-Fe associated coordination peaks at ~2.47 Å and ~2.57 Å for  $\alpha$ -Fe<sub>2</sub>O<sub>3</sub>,  $\gamma$ -Fe<sub>2</sub>O<sub>3</sub> and Fe<sub>3</sub>O<sub>4</sub> are marked by dashed vertical lines, respectively. Solid lines indicate evolution of Fe-Fe coordination peaks for 2-line ferrihydrite with hydrothermal treatment time. The dash-dot line indicates the position of Fe-O coordination peaks for the studied ferrihydrite samples.

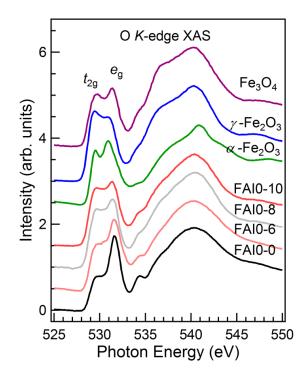
The influence of longer-range order on spectra. *L*-edge XAS spectra and *K*-edge XANES spectra for 2-line and 6-line ferrihydrite samples at room temperature are shown in Figure S5. XAS spectra at the *L*-edge for two samples were measured using linear polarized soft X-rays. As shown in Figure S5, the Fe *L*- and *K*-edge spectra for 2-line ferrihydrite resemble that of 6-line ferrihydrite, although 6-line ferrihydrite has longer-range order than that of 2-line ferrihydrite, which corresponds well with literature reports.<sup>2, 3</sup> It is not surprising that structural disorder has little influence on spectral shape in both *L*-edge XAS and *K*-edge XANES spectra because the spectra are dominated by quasi-atomic transitions, with minor modification from the solid state environment for the *L*-edge,<sup>4</sup> whereas for *K*-edge XANES spectra, the influence of structural order can be largely ignored, as the shape and intensity of main edge features are sensitive to short-range order (particularly the first coordination shell) around sites.



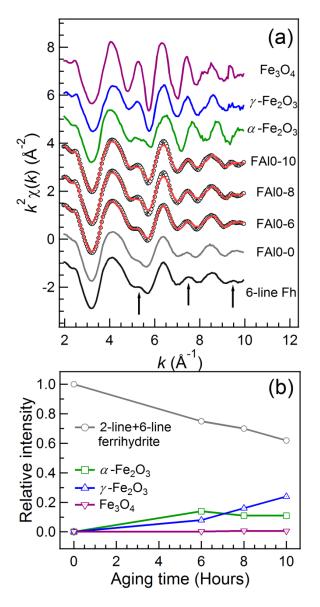
**Figure S5.** (a) *L*-edge XAS and (b) *K*-edge XANES spectra for 2-line and 6-line ferrihydrite. The difference spectrum is shown at the bottom of each panel.

Oxygen *K*-edge spectra for four studied samples and three reference phases of hematite, maghemite and magnetite are shown in **Figure S6**. The higher energy features between 533 eV and 545 eV are generally attributed to the transition from the O 1*s* to O 2*p* hybridized with the Fe

4*s* and 4*p* states. Pre-edge features in the 528~533 eV region are attributed to the transition from O 1*s* to O 2*p* states hybridized with Fe 3*d* states, which are split into  $t_{2g}$  and  $e_g$  states due to the oxygen octahedral crystal field.<sup>5,6</sup> The pre-edge features of the magnetite reference have higher positions and slightly worse resolution than that of hematite, which is in accordance with findings reported by Wu et al.<sup>7</sup> The formation of goethite is ruled out, otherwise a new feature located at the higher energy side of  $e_g$  resonance should emerge,<sup>8,9</sup> which is in good agreement with results from magnetic measurements.<sup>10</sup> It is worth noting that the measured XAS signal has a contribution from oxygen contamination due to air exposure, so that the accuracy of the intensity of each feature is unknown. However, this does not alter the main conclusion that  $\alpha$ - and  $\gamma$ - Fe<sub>2</sub>O<sub>3</sub> formed and contribute to evolution of the spectral profile at the O *K*-edge for hydrothermally aged ferrihydrite.



**Figure S6.** Oxygen *K*-edge XAS spectra for the four studied 2-line ferrihydrite samples and for three reference samples.



**Figure S7**. (a)  $k^2$ -weighted EXAFS spectra for four studied 2-line ferrihydrite samples, three iron oxide reference samples and a 6-line ferrihydrite reference. Arrows indicate oscillations for ferrihydrite, which are structural order dependent. Red solid lines represent best-fit curves for  $k = 2-10 \text{ Å}^{-1}$ . (b) Concentrations of different Fe oxides as a function of hydrothermal aging time.

In order to validate the linear combination fitting of K-edge XANES spectra, linear combination fitting of k spaced EXAFS spectra for studied samples was performed and results are shown as red curves in **Figure S7**(a). Considering the transformation of 6-line ferrihydrite

revealed by XRD and dependence of EXAFS spectra on structural order (arrows in **Figure** S7(a)), a 6-line ferrihydrite reference was selected. Compositions from the best fits are listed in **Table S1**. Evolution of each component vs hydrothermal aging time (**Figure S7**(b)) is consistent with linear combination fitting of *K*-edge XANES spectra, which confirms their validity.

**Table S1.** Contributions to  $k^2 \chi(k)$  EXAFS spectra from chemical species present for different hydrothermal aging times.

		$k^2 \chi(k)$ EXAFS spectra contribution (%)					1
XAS	Sample	FA10-0	6-line Fh	α-Fe <sub>2</sub> O <sub>3</sub>	γ-Fe <sub>2</sub> O <sub>3</sub>	Fe <sub>3</sub> O <sub>4</sub>	$-\chi^2$
<i>K</i> -edge $(k^2\chi(k))$	FA10-6	65	10	14	8	3	0.002
$(k \chi(k))$ EXAFS)	FA10-8	39	31	11	16	3	0.006
	FA10-10	35	27	11	25	2	0.005

## References

(1) Yang, T.-Y.; Wen, W.; Yin, G.-Z.; Li, X.-L.; Gao, M.; Gu, Y.-L.; Li, L.; Liu, Y.; Lin, H.; Zhang, X.-M.; Zhao, B.; Liu, T.-K.; Yang, Y.-G.; Li, Z.; Zhou, X.-T.; Gao, X.-Y. Introduction of the X-ray diffraction beamline of SSRF. *Nucl. Sci. Tech.* **2015**, 26 (2), 020101.

Masina, C. J.; Neethling, J. H.; Olivier, E. J.; Manzini, S.; Lodya, L.; Srot, V.; van Aken, P.
A. Structural and magnetic properties of ferrihydrite nanoparticles. *RSC Adv.* 2015, *5* (50), 39643-39650.

(3) Carta, D.; Casula, M. F.; Corrias, A.; Falqui, A.; Navarra, G.; Pinna, G. Structural and magnetic characterization of synthetic ferrihydrite nanoparticles. *Mater. Chem. Phys.* **2009**, *113* (1), 349-355.

(4) van der Laan, G.; Kirkman, I. W. The 2*p* absorption spectra of 3*d* transition metal compounds in tetrahedral and octahedral symmetry. *J. Phys.: Condens. Matter* **1992**, *4* (16), 4189.

(5) Signorini, L.; Pasquini, L.; Savini, L.; Carboni, R.; Boscherini, F.; Bonetti, E.; Giglia, A.; Pedio, M.; Mahne, N.; Nannarone, S. Size-dependent oxidation in iron/iron oxide core-shell nanoparticles. *Phys. Rev. B* **2003**, *68* (19), 195423.

(6) Chen, S.-Y.; Gloter, A.; Zobelli, A.; Wang, L.; Chen, C.-H.; Colliex, C. Electron energy loss spectroscopy and *ab initio* investigation of iron oxide nanomaterials grown by a hydrothermal process. *Phys. Rev. B* **2009**, *79* (10), 104103.

(7) Wu, Z. Y.; Gota, S.; Jollet, F.; Pollak, M.; Gautier-Soyer, M.; Natoli, C. R. Characterization of iron oxides by x-ray absorption at the oxygen K edge using a full multiple-scattering approach. *Phys. Rev. B* **1997**, *55* (4), 2570-2577.

(8) Gilbert, B.; Erbs, J. J.; Penn, R. L.; Petkov, V.; Spagnoli, D.; Waychunas, G. A. A disordered nanoparticle model for 6-line ferrihydrite. *Am. Mineral.* **2013**, *98* (8-9), 1465-1476.

(9) Brandt, F.; Schäfer, T.; Claret, F.; Bosbach, D. Heterogeneous formation of ferric oxide nanoparticles on chlorite surfaces studied by x-ray absorption spectromicroscopy (STXM). *Chem. Geol.* **2012**, *329*, 42-52.

(10) Liu, Q.; Barrón, V.; Torrent, J.; Eeckhout, S. G.; Deng, C. Magnetism of intermediate hydromaghemite in the transformation of 2-line ferrihydrite into hematite and its paleoenvironmental implications. *J. Geophys. Res.* **2008**, *113* (B1), B01103.