New Insights into the Molecular Structures, Compositions, and Cation Distributions in Synthetic and Natural Montmorillonite Clays

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Supporting Information

Cadars, S. et al., Molecular Structures of Synthetic and Natural Montmorillonite, Supporting Information

Atom	Core-states	Local channel	<i>r</i> _{loc} (a.u.)	r _{nonloc} (a.u.)	<i>r_{aug}</i> (a.u.)	pseudopotential projectors	PAW projectors
Н		р	0.8	0.8	0.6	2x2s	2x2s
Ο	1s	d	1.3	1.3	0.9	2x2s, 2x2p	2x2s, 2x2p
Na	1s	d	1.3	1.3	1.0	1x2s, 2x2p, 1x3s	1x2s, 2x2p, 1x3s
Si	1s, 2s,2p	d	1.8	1.8	1.3	2x3s, 2x3p	2x3s, 2x3p, 2x3d
Mg	1s, 2s	d	1.6	2.0	1.4	2x3s,1x2p, 2x3p	2x3s, 1x2p, 2x3p, 2x3d
Al	1s, 2s,2p	d	2.0	2.0	1.4	2x3s, 2x3p	2x3s, 2x3p, 2x3d

Table S1. Description of pseudopotentials used for planewave-based DFT calculations.

Pseudopotentials used for calculations on reference crystalline systems (see below)

В	1s	d	1.4	1.4	1.0	2x2s, 2x2p	2x2s, 2x2p
Ca	1s	f	1.6	2.0	1.4	1x3s, 2x3p, 1x4s	1x3s, 2x3p, 1x4s
Р	1s, 2s, 2p	d	1.8	1.8	1.3	2x3s, 2x3p	2x3s, 2x3p,2x3d
Κ	1s, 2s, 2p	d	1.8	1.8	1.6	1x3s, 2x3p, 2x4s	1x3s, 2x3p, 2x4s
Ge	1s, 2s, 2p, 3s, 3p, 3d	d	2.3	2.3	1.5	2x4s, 2x4p	2x4s, 2x4p
Y	1s, 2s, 2p, 3s, 3p, 3d	f	2.0	2.0	2.0	1x4s, 2x4p, 2x4d, 1x5s	1x4s, 2x4p, 2x4d, 1x5s

Where r_{loc} is the pseudisation radius for the local component of the pseudopotential, r_{nonloc} is the pseudisation radius for the non-local components of the pseudopotential, and r_{aug} is the pseudisation radius for the charge augmentation functions. The corresponding Materials Studio Castep on-the-fly strings used to generate these potentials are:

```
H 1|0.8|3.675|7.35|11.025|10UU(qc=6.4)[]
```

```
0 2 1.3 16.537 18.375 20.212 20UU: 21UU(qc=7.5)[]
```

```
Na 2 1.3 1.3 1 1.8 13.6 15.3 20U=-2.07:30U=-0.105:21U=-1.06U=+0.25[]
```

```
Mg 2 1.6 2 1.4 6 7 8 30NH:21U:31UU:32LGG(qc=4.5) []
```

```
Al 2 2 3.675 5.512 7.717 30UU: 31UU: 32LGG ]
```

```
Si 2|1.8|3.675|5.512|7.35|30UU:31UU:32LGG[]
```

```
B 2 1.4 9.187 11.025 13.965 20UU: 21UU (qc=5.5) ]
```

```
K 2 | 1.8 | 1.8 | 1.6 | 11 | 14.7 | 16.7 | 30U:40UU:31UU(qc=5.5) []
```

```
Ge 2 2.3 2.3 1.5 4.4 6 9 40U=-0.44U=+0.25:41U=-0.15U=+0.25 [] "
```

```
Y 3 2 2 2 8.5 10 11.1 40U:50U:41UU:42UU[]
```

```
Ca 3 1.6 2.0 1.4 7 9 10 30U:40U:31:32U=+0@+0.12U=+1.0@+0.12
```

```
The pseudopotential of Ca used the correction described by Profeta et al.<sup>1</sup>
```

Nucleus	Compond, formula	Site # Experiment shift (ppm		Reference	Calculated shielding (ppm)
	Brucite Mg(OH) ₂	1	0	2	30.8
	diaspore AlO(OH)	1	9.2	This work	19.8
	Boehmite AlO(OH)	1	7.2	This work	22.2
$^{1}\mathrm{H}$	Talc	1	0.7	3	30.6
	Pyrophyllite	1	2.2	3	28.2
	KH_2PO_4	1	14.9	4	12.2
	α -quartz SiO ₂	1	-107.4	5	429.6
	Cristoballite SiO ₂	1	-108.5	5	430.9
		3	-105		428.0
	albite NaAlSi ₃ O ₈	2	-97	6	419.0
		1	-93		414.4
	datolite CaBSiO ₄ (OH)	1	-83	5	404.4
	danburite CaB ₂ Si ₂ O ₈	1	-89	5	410.1
29~	Pyrophyllite Si ₄ Al ₂ O ₁₀ (OH) ₂	1	-94	6	416.7
²⁹ Si	Talc $Si_4Mg_3O_{10}(OH)_2$	1	-97	6	420.0
	Na ₂ SiO ₃	1	-76.8	5	396.9
	α -Na ₂ Si ₂ O ₅	1	-93.6	7	416.5
	R No Si O	1	-85.6	7	406.5
	β -Na ₂ Si ₂ O ₅	2	-87.5	/	408.0
	S No Si O	1	-90.6	8	411.7
	δ -Na ₂ Si ₂ O ₅	2	-90.2	δ	411.2
	Berlinite AlPO ₄	1	42.9	9	512
	YAlO ₃	1	10.7	10	545.31
	V ALO	1	77.5	10	480.04
	$Y_3Al_5O_{12}$	2	2.1	10	553.33
	Sillimonite Al SiC	1	63.9	11	491.72
²⁷ Al	Sillimanite Al ₂ SiO ₅	2	4.7		549.07
AI	Al ₂ Ge ₂ O ₇	1	36	11	516.15
		1	35	10	518.21
	Andalusite Al ₂ SiO ₅	2	10	12	543.34
	Low albite NaAlSi ₃ O ₈	1	62.7	13	486.69
	Albite NaAlSi ₃ O ₈	1	63	13	486.7

Table S2. Calculated shieldings (σ_{iso}) and experimental chemical shifts (δ_{iso}) of reference
systems of known crystal structure.

Calculations of shieldings for crystalline model systems of known structure and experimental shifts are used to accurately calculate the isotropic ²⁹Si chemical shifts (δ_{iso}) from calculated ²⁹Si shieldings (σ_{iso}). This procedure compensates for possible systematic errors of the DFT calculations. All calculations were conducted on structures previously optimized with fixed unit cell parameters. The series of compounds listed in Table S2 led to the following relationships: $\delta_{iso}(ppm) = -0.793 * \sigma_{iso} + 24.71$ for ¹H ; $\delta_{iso}(ppm) = -0.920 * \sigma_{iso} + 288.45$ for ²⁹Si ; and $\delta_{iso}(ppm) = -0.977 * \sigma_{iso} + 541.86$ for ²⁷Al.

Cadars, S. et al., Molecular Structures of Synthetic and Natural Montmorillonite, Supporting Information

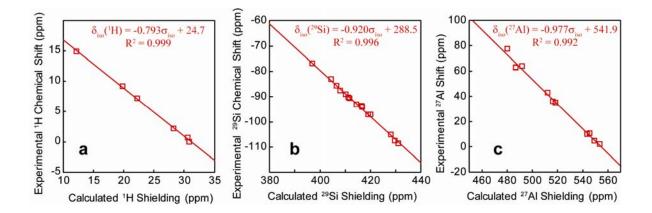


Figure S1. Correlation plots between experimental isotropic chemical shifts and isotropic shieldings calculated by DFT for the series of reference crystals of known structures listed in Table S2. (a), (b) and (c) correspond to ¹H, ²⁹Si, and ²⁷Al NMR data, respectively.

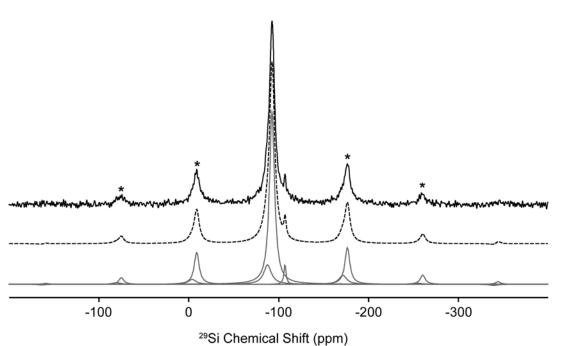


Figure S2. Solid-state NMR ²⁹Si echo-MAS spectrum of natural Na-Montmorillonite (same as in Figure 1c), collected at 8.8 T, 5 kHz MAS, using a short recycling delay of 200 ms (which considerably reduces the intensity of the slowly-relaxing quartz signal at -107 ppm). Spinning sidebands due to the effect of Fe³⁺ are marked with asterisks. The Q^3 (1Al) contribution is confirmed in both the isotropic region and the sidebands.

Cadars, S. et al., Molecular Structures of Synthetic and Natural Montmorillonite, Supporting Information

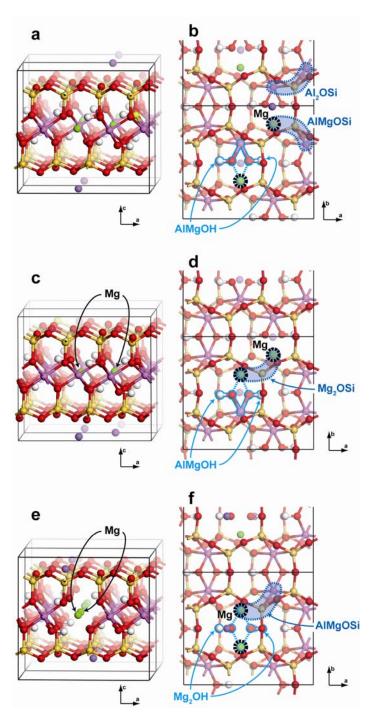


Figure S3. DFT-optimized models of Montmorillonite of composition $Na_2(Al^{(VI)}_6Mg_2)(Si_{16})$ $O_{40}(OH)_8$, *i.e.* with two Mg atoms in the octahedral layer. In (a) Mg atoms occupy nonadjacent sites, whereas they occupy adjacent octahedral sites forming either (b) Mg₂OSi or (c) Mg₂OH moieties.

Cadars, S. et al., Molecular Structures of Synthetic and Natural Montmorillonite, Supporting Information

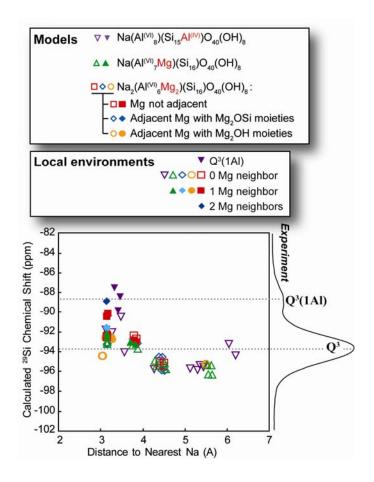


Figure S4. Plot of calculated ²⁹Si chemical shifts as a function of the distance to the nearest Na⁺ cation, using all Montmorillonite models shown in Figure 5 and Supporting Information, Figure S3. Symbol types and colour tones refer to different models, while open or filled symbols and variations of colour tones point to distinct types of local environments. There appears to be an effect of the Si-Na distance on the calculated ²⁹Si chemical shifts, especially between ca. 3.5 and 6 Å.

Cadars, S. et al., Molecular Structures of Synthetic and Natural Montmorillonite, Supporting Information

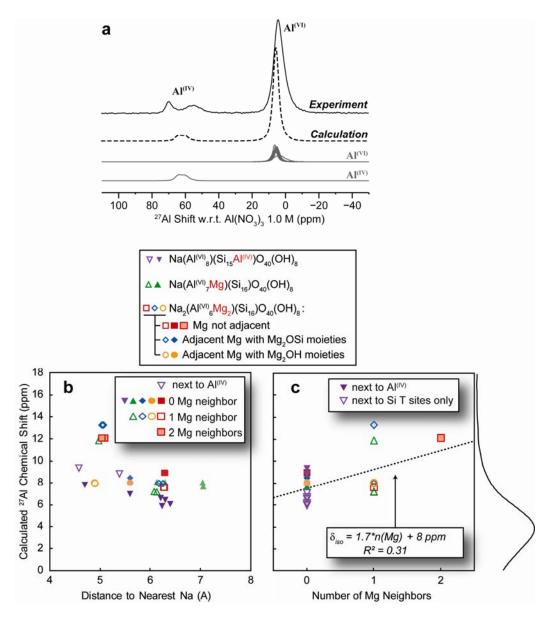


Figure S5. (a) Simulated ²⁷Al NMR spectrum (dashed line) obtained from DFT calculations of of NMR parameters, using the structural model composition $Na(Al^{(VI)}_{8})(Si_{15}Al^{(IV)})O_{40}(OH)_{8}$ (Figure 4a). The ²⁷Al MAS spectrum of Na-MMT (same as in Figure 2a) is shown on top for comparison (black solid line). Individual contributions from each site of the model to the calculated spectra are shown in grey lines. Arbitrary Lorentzian broadenings of 500 Hz were applied to each individual contribution. (b, c) Isotropic ²⁷Al chemical shifts calculated from DFT models shown in Figure 5 and Supporting information, Figure S3, and reported as functions of (b) the distance to the nearest Na atom, and (c) the number of Mg neighbors. Both parameters seem to contribute to some extent to the ²⁷Al shifts. We note that the model associated with yellow "o" is not representative of the actual material (absence of contribution from Mg₂OH moieties in ¹H NMR).

Cadars, S. et al., Molecular Structures of Synthetic and Natural Montmorillonite, Supporting Information

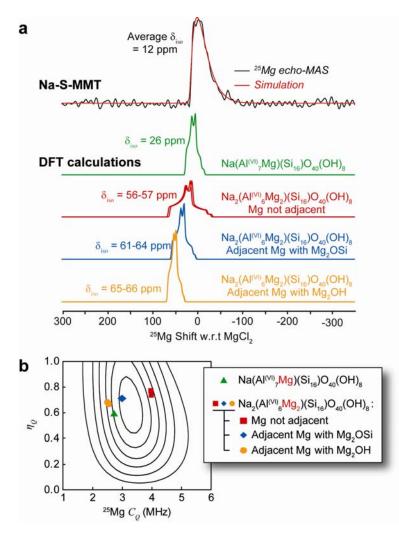


Figure S6. (a) ²⁵Mg MAS NMR spectrum of Na-S-MMT (in black, same as in Figure 3c) and corresponding simulation (in red) using the distribution of quadrupolar coupling parameters (C_Q, η_Q) plotted as contour levels in (b). The spectra represented in green, red, blue, and yellow in (a) correspond to the individual contributions calculated by DFT for the different Mg-containing models described above. The corresponding calculated quadrupolar coupling parameters are reported as symbols on top of the experimental distribution of quadrupolar coupling coupling parameters in (b).

Cadars, S. et al., Molecular Structures of Synthetic and Natural Montmorillonite, Supporting Information

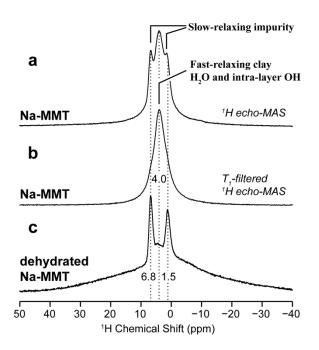


Figure S7. Solid-state ¹H echo-MAS NMR experiments performed on (a, b) natural montmorillonite (Na-MMT) and (c) dehydrated Na-MMT, at (a, b) 19.9 T and (c) 17.6 T, respectively, with a spinning frequency of 64 kHz. (a) Spectrum collected with a recycling delay of 10s. The contribution from the central peak at 4 ppm is underestimated due to its fast relaxation during the echo (two rotor periods). (b) Spectrum collected with a recycle delay of 0.1s and a large number of dummy scans to saturate the slow-relaxing components. (c) Spectrum collected with a recycling delay of 10s. The broad signal at the base is essentially due to rotor and probe background signals, which are not eliminated by the short echoes required to detect the fast relaxing clay signal.

Calculations of clay composition based on quantitative NMR data.

The composition of the tetrahedral layer may be calculated on the basis of ²⁹Si NMR data, assuming that the Al^(IV) atoms are completely isolated from each other in the tetrahedral layer. This is a reasonable hypothesis given the low amount of four-coordinated Al in the material and following the Lowenstein rule, which predicts the avoidance of Al^(IV)-O-Al^(IV) moieties and should at least partially apply in these materials.¹⁴ We used the following formula:

$$(\text{Si/Al}^{(\text{IV})})_{\text{NMR}} = \left(\sum_{n=0}^{3} I[Q^3(n\text{Al})]\right) / \left(\sum_{n=0}^{3} (n/3)I[Q^3(n\text{Al})]\right)$$

this yields a Si/Al ratio of 40 in the tetrahedral layer, corresponding to the composition $(Si_{7.80}Al_{0.20})$, calculated as:

$$n[\operatorname{Al}^{(IV)}] = \frac{n(\operatorname{Si} + \operatorname{Al}^{(IV)})}{1 + \operatorname{Si}/\operatorname{Al}^{(IV)}}$$

with $n(Si + Al^{(IV)}) = 8$, based on the theoretical composition of dioctahedral clays. The $Al^{(VI)}/Al^{(IV)}$ ratio is then used to derive the numbers of $Al^{(VI)}$, Mg, and inter-layer Na atoms.

Alternatively, the Mg and Al amounts in the octahedral layer may be obtained more directly from the relative amounts of AlMgOH and Al₂OH moieties using the expression:

$$n(Mg) = \frac{n(OH)}{2\left(\frac{n(Al_2OH)}{n(AlMgOH)} + 1\right)}$$

with n(OH) = 4 (based on the theoretical composition of dioctahedral clays), which integrates the absence of Mg₂OH moieties, as demonstrated by ¹H NMR. We note that this does not preclude the existence of adjacent Mg, since the formula still works if pairs of adjacent Mg atoms are present in the form of Mg₂OSi moieties, i.e. Mg atoms "connected" via an apical O. The numbers of Al^(IV), Si, and inter-layer Na atoms are then derived using the measured Al^(VI)/Al^(IV) ratio.

Cadars, S. et al., Molecular Structures of Synthetic and Natural Montmorillonite, Supporting Information

Trying to correct the ICP-EOS data to account for the amounts of Al and Si in the aluminosilicate impurity detected by NMR (*i.e.* 13% of total Al content, in the form of $Al^{(IV)}$ only, and 16% of the total Si content) is done as follows:

$$n[\mathrm{Al}^{(IV)}]_{\mathrm{MMT}} = n[\mathrm{Al}^{(IV)}]_{\mathrm{ICP-EOS}} \left(1 - \frac{n[\mathrm{Al}^{(IV)}]_{\mathrm{Impurity}}}{n[\mathrm{Al}^{(IV)}]_{\mathrm{Total}}}\right)$$

and

$$n(\text{Si})_{\text{MMT}} = n(\text{Si})_{\text{ICP-EOS}} \left(1 - \frac{n(\text{Si})_{\text{Impurity}}}{n(Si)_{\text{Total}}}\right)$$

We obtain a corrected $[Si/Al^{(IV)}]_{MMT}$ ratio of 85, which is not very different from the original $(Si/Al^{(IV)})_{ICP}$ ratio and still far from the ratios (between 30 and 40) calculated from NMR data.

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